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MESOZOIC AND CENOZOIC PLUTONIC DEVELOPMENT IN THE ANDES OF CENTRAL CHILE (30°30'-32°30')

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Mesozoic and Cenozoic granitoids of this segment of the Andes (Fig.1) which have been studied by RIVANO et al. (1985) are grouped here in three NS- trending belts; western (WB), central (CB) and eastern (EB). The most important features of the belt components are summarized in Table 1. The WB is formed by the Mincha and Illapel superunits; the CB includes the Cogotí superunit and the San Lorenzo unit and the EB is occupied by the Rio Grande and Rio Chicharra superunits.

The country rocks adjacent to the western flank of the WB comprise Paleozoic lowgrade metasediments and metavolcanic rocks, together with Lower Mesozoic sediments and siliceous volcanics. Various roof-pendants are recognised in the central and eastern part of the WB. They consist of volcanogenic sequences and less frequent clastic sediments and limestones assigned to Lower Cretaceous formations. Plutons of the CB are roofed by gently folded Upper Cretaceous volcaniclastic sequences. The envelope of the EB comprises several sedimentary and volcanic formations with ages decreasing eastwards from Upper Jurassic to Paleogene. There is no evidence of deposition of younger cover rocks over the plutonic belts.

K-Ar ages show discrete ranges for each of the belts (Table 1) with a pronounced eastward migration of magmatism with time: WB (Jurassic to Late Cretaceous), CB (Early Tertiary), EB (Late Tertiary). Some units within the WB have now also been dated by the Rb-Sr whole-rock isochron method, which suggests some N-S migration within the Puerto Oscuro unit. The jumps in the sites of the magmatic belts correspond to essentially non-magmatic intervals (85-70 Ma) and (38-36 Ma) may relate to periods of subduction erosion or changes in the dip angle of the subducted lithosphere. Particularly rapid migration corresponds to very rapid oceanic spreading at 110-85 Ma (18 cm/yr; LARSON and PITMAN, 1972) and fragmentation of the Farallon plate to the present day configuration at 25 Ma ago (WORTEL, 1984).

Within each superunit the chemistry shows typical major element oxide variations with trends of increasing K_2O and decreasing Al_2O_3 , CaO and FeO with respect to increasing SiO_2 , but Na_2O is rather constant. TiO₂ and P_2O_5 increase to a maximum in rocks of intermediate composition (47-60% SiO_2) and thereafter decrease slowly. There are few distinctive differences between the chemical compositions of different superunits: high Al_2O_3 , CaO, Na_2O and P_2O_5 and low K_2O and Rb content in the Limahuida granitoids are characteristic of a trondhjemitic association (BARKER and ARTH 1976). The transition metal contents of the intermediate granitoids (56-63% SiO_2) are comparable with those of oceanic island-arc andesite (BAILEY 1981). Very low initial ⁸⁷ Sr/⁸⁶ Sr

TABLE 1

SUMMARY OF PETROGRAPHIC AND CHEMICAL FEATURES FOR MESOZOIC AND CENOZOIC GRANITOU SUPERUNITS BASED ON ANALYSES FROM RIVANO AND SEPULVEDA (IN PREP.)

BELT	SUPERUNIT	UNIT	PETROGRAPHIC & TEXTURAL FEATURES	AREA PROPORTION+	CHEMICAL FEATURES
WESTERN BELT	NINCHA 200-143 Na; RD-56 191-138 Na; K-Ar	Hillahue	Medium-grained monzogranite and sieno- granite; Ci \leq 10.	~ 158	73-77% \$102 (75.0); 12.4-13.8 (12.9); 0.01-1.34% CaO (0.60) Na20: 0.78-1.21; A/CNK: 0.59- K/Rb: (189); Rb/Sr: (3.0).
		Tranquilla	Slenogranite;(; < 10. Equigranular, fine grain size, porphyrysk, graphic.	~ 5t	73-76% SIO ₂ (74.9); 12.5-13.6 (13.0); 0.1-0.92% CaO (0.57); 4.4% Ma ₂ O (4.0); 3.4-5.1% K ₂ K ₂ O/Ma ₂ O: 0.83-1.36; A/CMK ; 1.16 K/Rb: (213); Rb/Sr (3.3)
		Püertö Öscura	Two-pyroxene diorite and monzodiorite. Minor olivine gabbro and quartz diori- te. CI : 30-50, idiomorphic, medium grain size. Local mineral orientation.	~ 358	47-57% SIO ₂ (53.7); 16.1-19.7 (18.1) 0.9-3.8% Ne ₂ O (2.9); 2.72% K ₂ O (1.3) K ₂ O/Ne ₂ O: 6.4 K/Rb: (197); Rb/Sr: (0.13); ppm Cr in rocks with > 50% Si 21 ppm Hi in rocks with > 50%
		Cavî lotên	Nb + Bt tonalite and granodiorite. Cl :10- 30, medium grain size, isotropic. Mafic inclusions.	~ 452	51-67% Si02(64.3); 15.5-17.8 (16.1) 3.7-5.7% CaO (4.7); 3. Na2O (3.7); 1.9-3.0% K2O (2.) A / CHK : 0.92-0.98; K2O/46 0.90 K/Rb :(213); Rb/Sr: (8.)
	111.00EL 134-86 No 134-86 No	Chalinga	Hb + Bt + Pa tonalite and quartz diorite.	~ 902	53-648 SIO2 (60.7); 7.3-2.48
			ç; 10 − 30, medium grain size.		(4.98) 3.4-6.5% Na ₂ 0 (4.1); 1 K ₂ 0 (2.3); A/CHK < 1.0; K ₂ M 0.35-1.01; K/Rb: (315); Nb/S
		L i mahu i da	Leucodiorita, trondjhemite with Cpx and Hb. Medium grain size.	∿ 10 t	50-70% S102 (61.7); 2.6-10.3 (7.8) 3.6-6.1% Mago (4.9); 8, Kg0 (0.4); high A1203 and Ag A/CHK: 0.68-1.09 Kg0/Mago: 8, K/Rb: (209) Rb/Sr: (0.02).
	006011 67-38 Na K-Ar	Fredes	Hb + Cpx, diorite, quartz monzodiori-	~ 907	50-692 510, (59.4); 144-1958
		rredes	te and granodiorite. Minor oliving	8 800 8	(17.0); 2.4-9.72 (a0 (5.6) 3.
CENTRAL BELT	= 2		gabbro. Fine to medium grain size		5.2% Na20 (3.9); 0.3 - 4.4% K (2.2); A/CNK: 0.78-1.03; K20/
	C060 67-3		seldom porphyrytic, Cl: 5-35.		0.22165; K/Rb: (2.61); Rb/1 (0.56).
	a a	Nogalada	Mainly fine-grained laucogranite. CI 5	< 10%	63-771 SIO2 (72.7): 12.2-15.6 A1203 (13.6) 41-0.632 CoO (1. 3.3-4.32 Ne20 (3.7): 3.1-5.5
		* * *			(4.6) K ₂ 0/Ma ₂ 0: 0.72-1.52; A 0.88-1.15 K/Rb: (278); Rb/Sr (2.04).
		San Lorenzo	Fine grained diorite, andesite porphy-		49-572 \$102(53.0); 4.7-8.52
		(65 ma)	ry, with hornblende and clinopyroxame.		(7.2); 2.7-5.62 ма ₂ 0 (3.9); (2.03 к ₂ 0 (1.4) к ₂ 0/ма ₂ 0: 0.1 А/сык: 0.77-1.0; к/Rb: (311); Sr (0.07).
EASTERN BELT	NIO GRANDE 36-28 Ma.K-Ar	Rio Tascadero	Hb + Bt monzonite and monzodiorite. Cl < 20.	< 20%	58-66% 5102 (62.4); 2.7-5.8% (4.6); 4.0-4.6% Na20 (4.2) 1 4.1% Kg0 (2.7); A/CNK: 0.83-
			e e		K ₂ O/Ha ₂ O: 0.44-1.0; K/Rb: (2 Rb/Sr: (0.33).
		Rio Las Cueves	Leucononzogranite and leucogranodiori- te. Medium grain size.	802	68-75% SiO ₂ (72.4); 1.0-3.5% (1.89) 3.6-4.2% Ma ₂ O (3.9) 3 4.9% K ₂ O (4.2); 12.6-15.8% A (13.9): A/CNK: 0.93-1.0 K ₂ O/ 0.81-1.16.
	CHICHARAA	Rio Cerro Blanco	Monzodiorite with variable amount of	902	\$1-672 \$102 (58.6); 3.2-9.5
			hornblende, clinopyroxene and biotite. Subidiomorphic, medium grain size and Cl: 15-35.	2 B	(6.2); 2.8-5.7% ма ₂ 0 (4.4) 4.3% К ₂ 0 (2.0); А/СИК: 0.76 К ₂ 0/Ма ₂ 0: 0.20-1.1; К/Rb: В (5.2) (0.20)
	AIO CHI 17- 0 M	Portezuelo del Azufre	Dacite porphyry. Felsitic and Inter- granular textures.	~ 102	Ab/Sr: (0.25).

Between 31° to 32°5. Numbers in bracket: average: A/CNK. Molar ratio Al₂0₃/(LaD + Ka₂0 + K₂0). Rb-Sr dates were obtained at the British Goo cal Survey. Most of the K-Ar ages were determined at the Servicio Nacional de Geología y Minería (Alvano at W. 1988)



Fig. 1. Mesozoic and Cenozoic plutonic rocks distribution in the Andes of Central Chile (30° 32'-32° 30): units and super units.

- 1) Tranquilla and Millahue Units
- 2) Puerto Oscuro and Cavilolén Units
- 3) Chalinga Unit
- 4) Limahuida Unit

- 5) San Lorenzo Unit
- 6) Cogotí Super Unit
- 7) Rio Grande Super Unit
- 8) Rio Chicharra Super Unit
- For more detailed information see Table
- 9) K-Ar ages in Ma. Errors quoted at the 2 sigma level.
- 10) Whole rock Rb-Sr isochron ages in Ma. Errors quoted at the 2 sigma level.
- 11) Average of 8 K-Ar ages in Ma.
- 12) Intrusive and Tectonic contacts of plutonic rocks.

rations for the WB plutonic rocks (mostly ca.0.7035) confirm direct derivation of the parent magmas from the upper mantle with virtually no continental crustal involvement. This distinguishes the Mesozoic-Cenozoic granitoids from those of the Paleozoic belts of Chile and signifies a fundamental change in the conditions of magma generation in late Triassic times.

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